Thermochemical models of molten silicates and silicate vapor

atmospheres on hot rocky exoplanets

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Outline

Thermochemical modeling of molten silicates Thermodynamics of Solutions Ideal Mixing of Complex Components Model Some examples with the MAGMA code Applications to hot rocky planets

Thermochemical modeling of molten silicates

In general, activity a_i and mole fraction X_i are related via $a_i = \gamma_i X_i$ with $\gamma_i = 1$ for ideal solutions and $\gamma_i \neq 1$ for non-ideal solutions

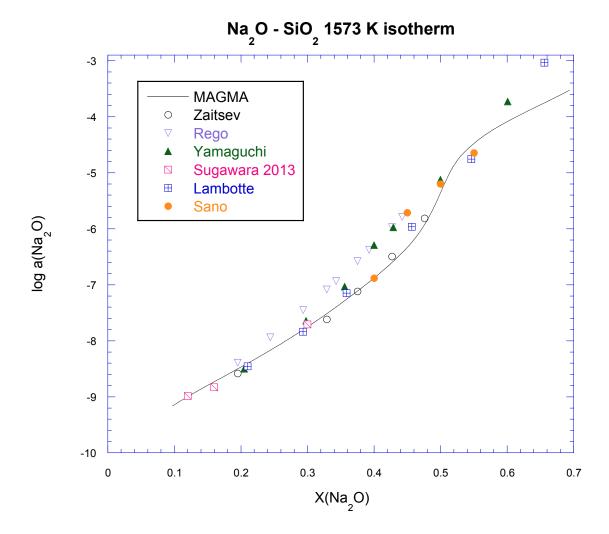


Figure shows activity << mole fraction for Na₂O

We want to compute a = f(T,P,X) for multicomponent silicate melts.

Then we can compute all other properties of the silicate melts.

Define some concepts and nomenclature for thermodynamics of solutions before proceeding

Thermodynamics of solutions

Mole Fraction Notation

We express concentration as mole fraction X_i , e.g., for the Na₂O –

SiO₂ binary solution, the mole fraction of Na₂O is

$$X_{Na_2O} = \frac{n_{Na_2O}}{n_{Na_2O} + n_{SiO_2}}$$

where n is the number of moles of Na₂O or SiO₂, respectively.

The sum of mole fractions for a solution is unity. For example,

considering molten soda-lime glass (Na₂O – CaO – SiO₂)

$$X_{Na_2O} + X_{CaO} + X_{SiO_2} = 1$$

Typical Na₂O/SiO₂ molar ratios in planetary materials:

Object	Na ₂ O/SiO ₂ molar ratio
Bulk silicate Moon	0.0010
Bulk silicate Earth	0.0076
Earth's crust	0.052
CI Chondrites (= solar ratio)	0.029

The SiO₂-rich end of $Na_2O - SiO_2$ melts is most relevant for us.

Thermodynamics of solutions II.

Gibbs Energy of Solution

Two or more components form solutions because it is energetically favorable to do so – the solution has a lower Gibbs energy G than the physical mixture of the two (or more) components.

$$\Delta G^M = G_{solution} - G_{mixture}$$

We call ΔG^{M} the Gibbs energy of mixing. For the Na₂O – SiO₂ binary,

$$\Delta G^{M} = n_{Na_{2}O} \left(\bar{G}_{Na_{2}O} - G^{o}_{m,Na_{2}O} \right) + n_{SiO_{2}} \left(\bar{G}_{SiO_{2}} - G^{o}_{m,SiO_{2}} \right)$$
$$\Delta G^{M} = n_{Na_{2}O} \left(RT \ln a_{Na_{2}O} \right) + n_{SiO_{2}} \left(RT \ln a_{SiO_{2}} \right)$$

The terms in the equations are the number of moles (n_i) of each component, the partial molal Gibbs energy of each component in the solution (e.g., \bar{G}_{Na_2O}), the molar Gibbs energy of each pure liquid oxide (e.g., G_{m,Na_2O}^o) at the same temperature, and the activity (a_{Na_2O} , a_{SiO_2}) of each component (see Solution Thermodynamics slide IV). All the other thermodynamic properties of solutions can be derived by differentiation of ΔG^M the Gibbs energy of mixing using the fundamental equation dG = VdP – SdT and the Gibbs – Helmholtz equation G = H – TS.

Thermodynamics of solutions III.

Other Properties Derived From ΔG^{M}

For example, the entropy of mixing is given by the partial derivative at constant pressure (P) and composition (X)

$$\Delta S^M = -\left(\frac{\partial \Delta G^M}{\partial T}\right)_{P,X}$$

The volume of mixing is given by

$$\Delta V^M = \left(\frac{\partial \Delta G^M}{\partial P}\right)_{T,X}$$

The enthalpy (or heat) of mixing is given by

$$\Delta H^M = \left(\frac{\partial \Delta G^M / T}{\partial (1/T)}\right)_{P,X}$$

The constant pressure heat capacity $C_P(\partial H/\partial T)$ of mixing is given by

$$\Delta C_P^M = -T \left(\frac{\partial^2 \Delta G^M}{\partial T^2} \right)_{P,X}$$

Thermodynamics of solutions IV.

Definition of Thermodynamic Activity

Thermodynamic activity (a) – henceforth activity – G.N. Lewis 1907

$$a = \frac{f}{f^*} \cong \frac{p}{p^*}$$

The activity of a material is approximately equal to the ratio of the partial pressure (p) of the material at some P,T,X and the saturation vapor pressure (p^{*}) of the pure material at the same total pressure (P) and temperature (T). The activity is exactly equal to the fugacity ratio where $f = \Gamma p$ with $\Gamma = 1$ for ideal and $\Gamma \neq 1$ for non-ideal gases. In the absence of PVT data for many high temperature gases, we assume $\Gamma = 1$ for our discussion and define activity as

$$a = \frac{p}{p^*}$$

Example: At 1600 C (1873 K) the vapor pressure of pure liquid Fe is 0.0572 mm Hg and the Fe vapor pressure over molten Fe – Ni alloy $(Fe_{0.9}Ni_{0.10})$ is 0.05124 mm Hg. Calculate the activity of pure liquid Fe and Fe in the molten alloy from these data.

$$a_{Fe(liq)} = \frac{0.0572}{0.0572} = 1$$
 and $a_{Fe(alloy)} = \frac{0.05124}{0.0572} = 0.90$

Thermodynamics of solutions V.

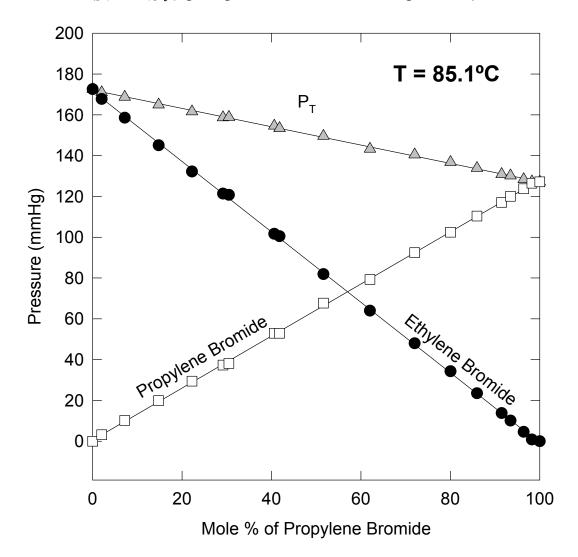
Ideal Solutions

These follow Raoult's Law at all concentrations:

$$\boldsymbol{a_i} = \boldsymbol{X_i} = \frac{f_i}{f_i^o} \cong \frac{p_i}{p_i^o}$$

Ethylene and propylene bromide $(C_2H_4Br_2 - C_3H_6Br_2)$ form an ideal

solution ($p_i = X_i p_i^o$) giving a characteristic straight-line plot:



Thermodynamics of solutions VI.

Enthalpy of Mixing

Ideal solutions also are athermal

$$\Delta H^M = 0 = \left(\frac{\partial \Delta G^M / T}{\partial (1/T)}\right)_{P,X}$$

This is easy to understand if we use the definition of an ideal solution

and substitute into our expression for the Gibbs energy of mixing.

Any two components A and B in an ideal solution $(a_i = X_i)$

$$\Delta G^{M} = n_{A}(RTlna_{A}) + n_{B}(RTlna_{B})$$
$$\Delta G^{M} = n_{A}(RTlnX_{A}) + n_{B}(RTlnX_{B})$$

The ideal gas constant R, number of moles, and mole fractions are constant, hence

$$\left(\frac{\partial \Delta G^M/T}{\partial (1/T)}\right)_{P,X} = \frac{\partial}{\partial (1/T)} R[n_A ln X_A + n_B ln X_B] = 0$$

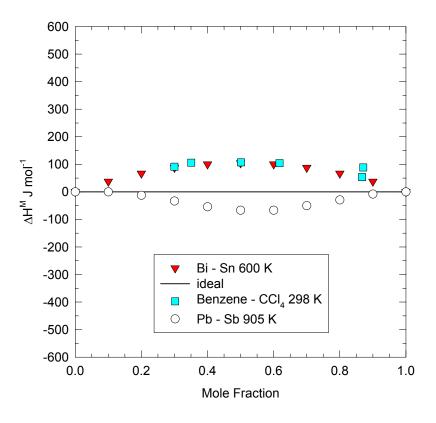
Conversely non-ideal solutions have $\Delta H^M \neq 0$ because $\gamma_i \neq 1$

$$\left(\frac{\partial \Delta G^M/T}{\partial (1/T)}\right)_{P,X} = \frac{\partial}{\partial (1/T)} R[n_A ln\gamma_A + n_B ln\gamma_B] \neq 0$$

The next page shows ΔH^M for ideal and non-ideal solutions.

The large negative ΔH^M values for Na₂O – SiO₂ occur because

activity coefficients << 1 for Na₂O



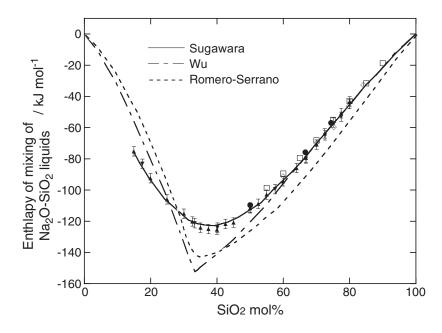


Fig. 2. Enthalpy of mixing of Na₂O–SiO₂ liquids: \bullet , present study (DSC); \blacktriangle (glass) and \checkmark (gel), Fan [10] (Transpose-temperature drop calorimetry); \Box , Hovis et al. [35] (HF solution calorimetry); \diamond , Rogez and Mathiue [51] (Lead borate solution calorimetry).

Ideal Mixing of Complex Components (IMCC) Model

Developed by Hastie & colleagues at NIST(NBS) in 1980s

Very negative ΔH^{M} values (e.g., Na₂O – SiO₂) show a_{oxide} << X_{oxide}

The IMCC model computes

$$a_{oxide} = X^*_{oxide}$$

The X_{oxide}^* is the mole fraction of unbound oxide in the melt The activity coefficient is calculated from

$$\gamma_{oxide} = \frac{X^*}{X^T}$$

The X^T is the total mole fraction of the oxide added to the melt.

Combination of these two equations is simply

$$a_{oxide} = \gamma_{oxide} X^T$$

The change from X^T to X^*_{oxide} is computed using chemical equilibria

between fictive components (aka pseudo-species) in the melt.

 K_2SiO_3 (liq), $K_2Si_2O_5$ (liq), and $K_2Si_4O_9$ (liq) used for $K_2O - SiO_2$ melts

IMCC model – precursor of models by Allendorf, Spear & colleagues Does not depend on actual (poorly known) speciation in the melts Good agreement with experimental data (shown later) Can be modified for oxides with $\gamma_{oxide} > 1$

MAGMA Code Example for K₂O – SiO₂ melts

1. $K_2O(liq) + SiO_2(liq) = K_2SiO_3(liq)$

$$K_1 = \frac{a_{K_2 S i O_3}}{a_{K_2 O} a_{S i O_2}}$$

2. $K_2O(liq) + 2 SiO_2(liq) = K_2Si_2O_5(liq)$

$$K_2 = \frac{a_{K_2 S i_2 O_5}}{a_{K_2 O} a_{S i O_2}^2}$$

3. $K_2O(Iiq) + 4 SiO_2(Iiq) = K_2Si_4O_9(Iiq)$

$$K_3 = \frac{a_{K_2 S i_4 O_9}}{a_{K_2 O} a_{S i O_2}^4}$$

The equilibrium constants are calculated from literature data

The activity coefficient for K₂O is given by

$$\gamma_{K_2O} = \frac{a_{K_2O}}{a_{K_2SiO_3} + a_{K_2Si_2O_5} + a_{K_2Si_4O_9}}$$

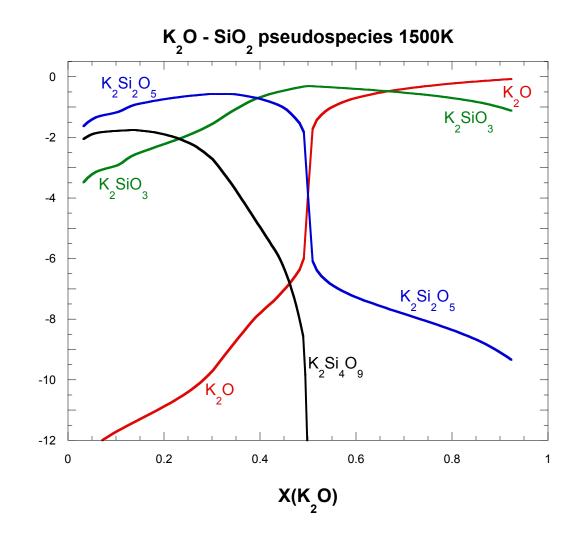
The activity of K_2O calculated from the reactions above and from the vaporization reaction below must be equal.

2 K₂O (liq) = 4 K (gas) + 2 O₂ (gas)

$$K_{vap} = \frac{P_K^4 P_{O_2}^2}{a_{K_2O}^2}$$

Gas phase equilibrium chemistry including thermal ionization, also computed, e.g., K, K₂, K⁺, e⁻, KO, K₂O for potassium

See figure showing pseudo-species, 1500 K isotherm



log Activity

MAGMA Code

Includes SiO₂, MgO, FeO, CaO, Na₂O, K₂O, TiO₂, and Al₂O₃

Most abundant oxides in silicates on Earth and other rocky planets

Consequence of two factors

Table 1

- solar abundances of the elements

- geochemical affinity (atmophile, chalcophile, lithophile, siderophile)

Species included in MAGMA code calculations Melt species Vapor species SiO₂,MgO, FeO, CaO, Al₂O₃, TiO₂, $0, 0_{2}$ Si, SiO, SiO₂ Na₂O, K₂O Mg, MgO MgSiO₃, Mg₂SiO₄, MgAl₂O₄, MgTiO₃, MgTi₂O₅, Mg₂TiO₄,Mg₂Al₄Si₅O₁₈ Fe, FeO FeTiO₃, Fe₂SiO₄, FeAl₂O₄ Ca, CaO CaAl₂O₄, CaAl₄O₇, Ca₁₂Al₁₄O₃₃, CaMgSi₂O₆, Al, AlO, AlO₂, Ca₂MgSi₂O₇, Ca₂Al₂SiO₇, CaAl₂Si₂O₈, CaTiO₃, Al_2O, Al_2O_2 Ca₂SiO₄, CaTiSiO₅, CaAl₁₂O₁₉, CaSiO₃ Ti, TiO, TiO₂ Al₆Si₂O₁₃ Na, Na₂, NaO, Na₂SiO₃, Na₂Si₂O₅, NaAlSiO₄, NaAlSi₃O₈, Na_2O, Na^+ NaAlO₂,Na₂TiO₃, NaAlSi₂O₆ $K, K_2, KO,$ $K_{2}O, K^{+}$ $K_2SiO_3, K_2Si_2O_5, K_2Si_4O_9, KAlSiO_4,$ KAlSi₃O₈,KAlO₂, KAlSi₂O₆, KCaAlSi₂O₇ e⁻

Look at some examples of code calculations vs.

- experimental data

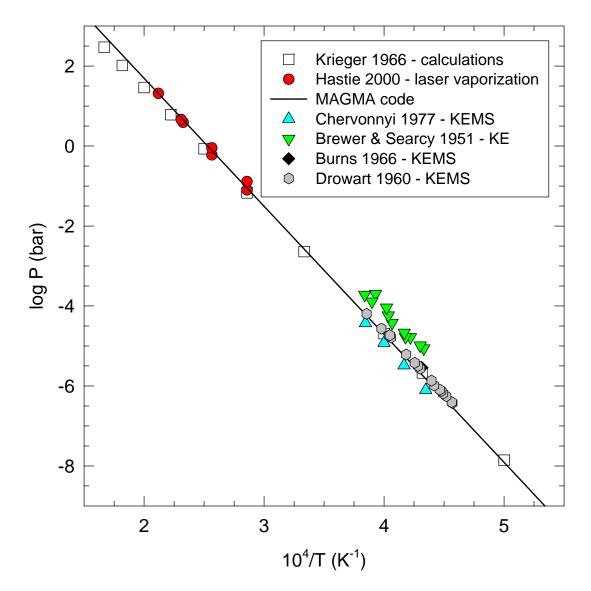
- theoretical calculations by other groups

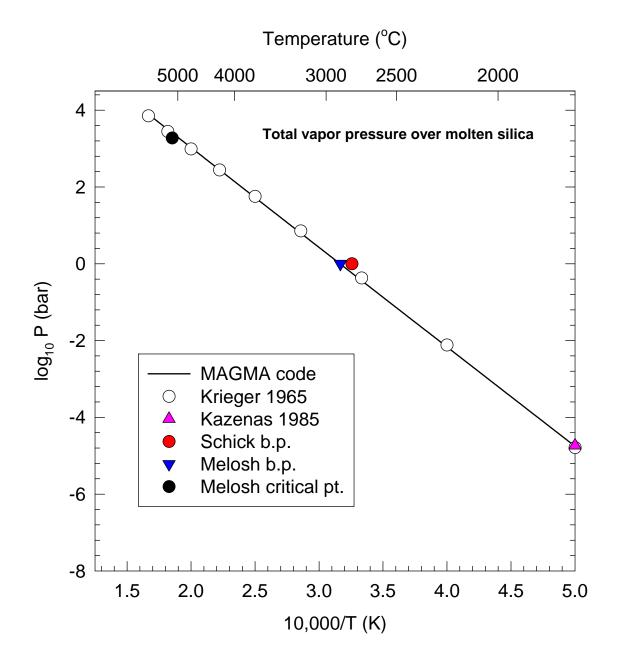
Vapor pressures of pure oxides

 $Na_2O - SiO_2$ and $K_2O - SiO_2$ binary systems

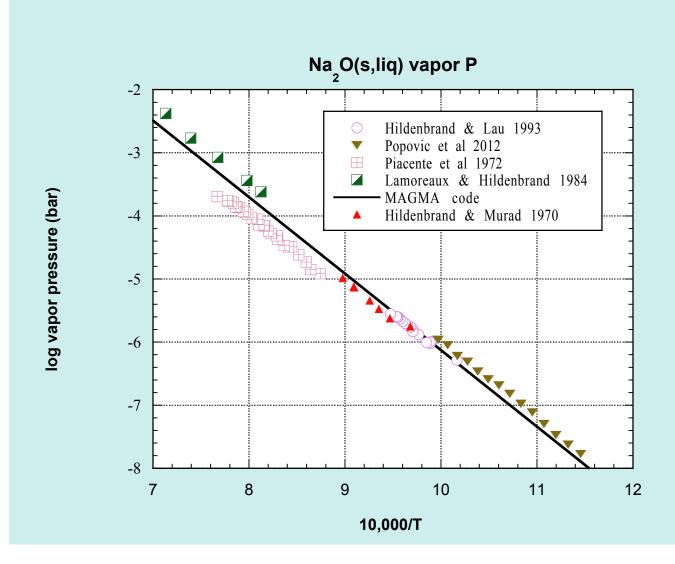
Multicomponent systems

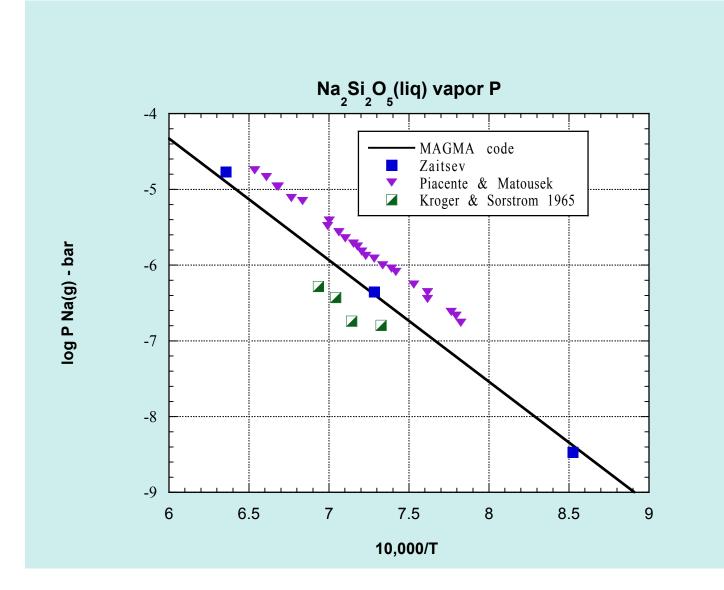
Alumina Vapor Pressure

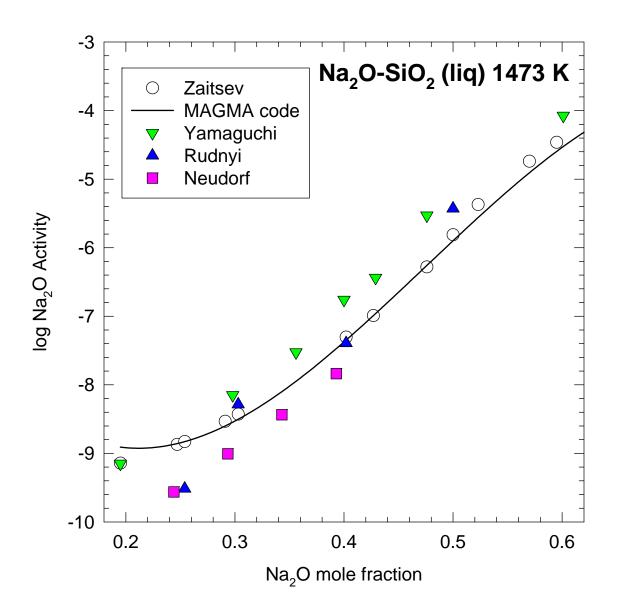




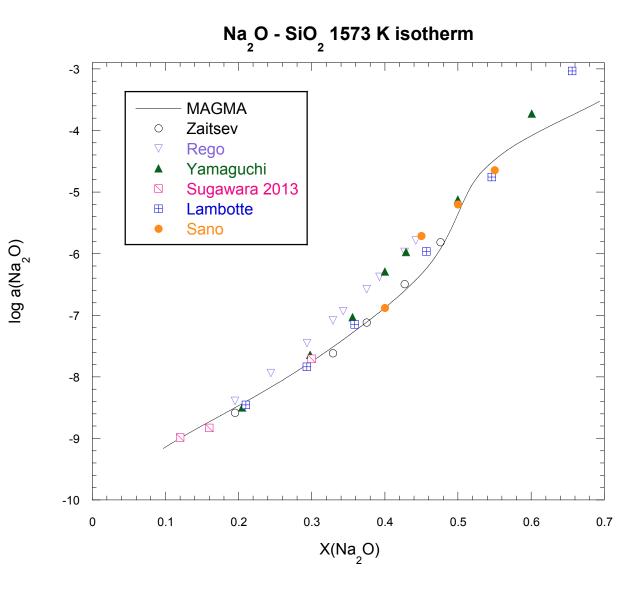
SiO2vap.SPW

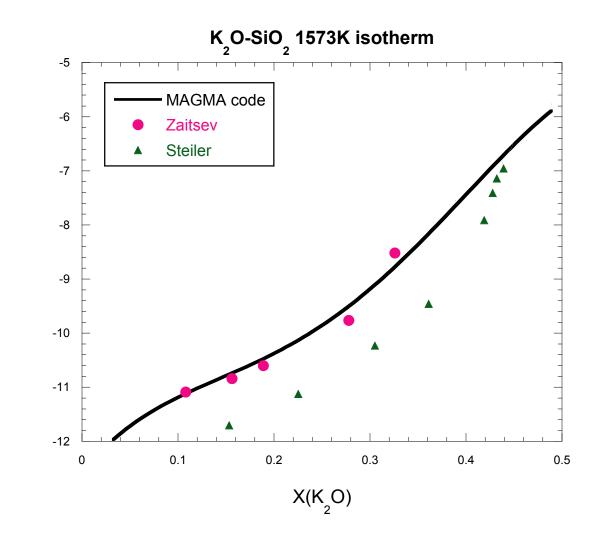




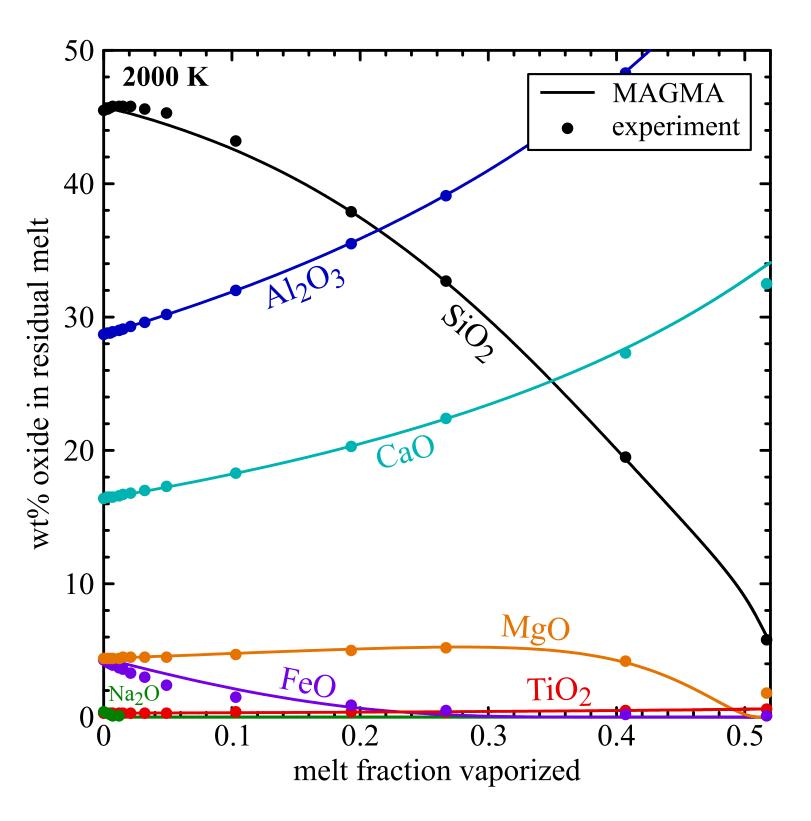


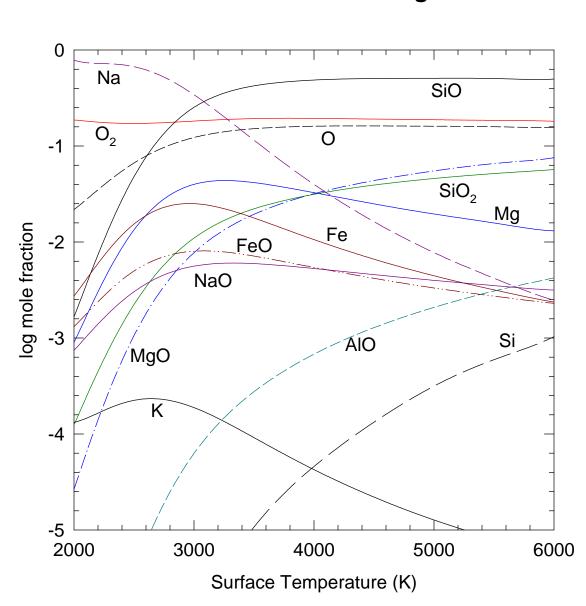
Na2OSiO2.SPW



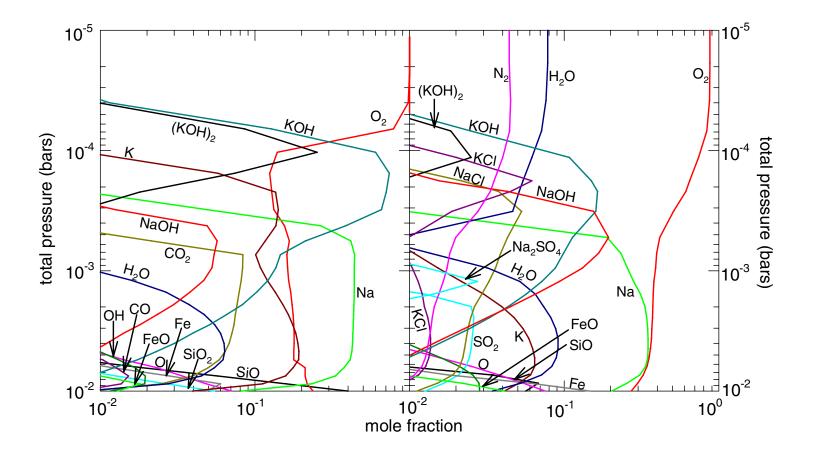


log a(K₂O,liq)





Composition of saturated vapor Bulk silicate Earth magma



CoRoT-7b dry adiabatic P-T profiles continental crust (left), BSE (right) surface T 2500 K, surface P 0.01 bar

Sodium oxide.

Sublimation and vaporization of Na₂O(s, liq) occur congruently and proceed mainly via the net reaction

 $2Na_2O(s, liq) = 4Na(g) + O_2(g)$

Only minor amounts of NaO, Na₂O, Na₂O₂, NaO₂ in the vapor.

The calculated vapor pressures from the MAGMA code for $Na_2O(s)$ from 950 – 1405 K and for Na_2O (liq) from 1405 – 2000 K are

Na₂O (solid) 950 – 1405 K: $log_{10}P$ (bar) = 5.995 – 12,124.1/T Na₂O (liquid) 1405 – 2200 K: $log_{10}P$ (bar) = 5.091 – 10,854.4/T

These two equations give a triple point pressure of ~ 0.0023 bar and a one bar boiling point of 2132 K. At the boiling point the saturated vapor is ~ 78% Na(g), 19% O₂(g), 1.5% NaO(g), 1.0% Na₂O(g), and 0.5% all other gases (O, O⁺, Na⁺, e⁻, Na₂).

Water Vapor and Hydroxide Gases

In the absence of water the vaporization of alkali metals from silicate melts proceeds via net thermochemical reactions exemplified by

$$2Na_2O$$
 (melt) = $4Na(g) + O_2(g)$

Water. Water vapor increases the amount of alkali metals in the gas phase over silicate melts via net thermochemical reactions such as $H_2O(g) + Na_2O$ (melt) = 2NaOH(g)

$$K_{OH} = \frac{P_{NaOH}^2}{P_{H_2O} a_{Na_2O}}$$

The ratio of the NaOH and Na gas pressures over a melt is

$$\frac{P_{NaOH}}{P_{Na}} = \left[\frac{K_{OH}P_{H_2O}}{K_{vap}f_{O_2}^{-\frac{1}{2}}}\right]^{\frac{1}{2}} > 1$$

The NaOH/Na molar ratio is > 1 unless P_{H_2O} is sufficiently small that

$$K_{vap} > K_{OH} P_{H_2O} f_{O_2}^{\frac{1}{2}}$$

Argent et al (1979) calculated Na gas speciation above a float glass melt (by moles: 12% Na₂O, 10% CaO, 6% MgO, 72% SiO₂) and found NaOH is the predominant Na gas at 1300 – 1800 K, in a gas containing (by volume) 0.5-4% O₂, 10-20% water vapor.

MAGMA code and hand calculations at 1600 K give

NaOH/Na molar ratio ~ 5.4

Future plans

- Complete comparison of K₂O SiO₂ binary with experimental data & other theoretical models
- 2. Include newer thermodynamic data for other key binaries
 - a. $Na_2O Al_2O_3$ and $K_2O Al_2O_3$ binaries
 - b. $CaO AI_2O_3$ and MgO SiO₂ binaries
 - c. MgO SiO₂ and CaO SiO₂ binaries
- 3. Update $M_2O SiO_2 AI_2O_3$ ternary
- 4. Add S- and halogen-bearing gases into code
 - Can be major species for alkalis Schaefer et al (2012)
- 5. Add H-bearing gases into code water, hydroxides, hydrides
 - 1st approximation no water solubility in melts T > 2000 K
 - No water solubility data in this T-range

Revisions motivated by newer thermodynamic data for liquid oxides and silicates and the need to compare different new datasets for the liquid oxides and silicates (e.g., geological vs. materials science)

Some Suggestions for Experimental Studies

Vaporization of peridotite rock – major constituent of Earth's mantle (Using Earth as a starting point to study extrasolar rocky planets, could also use the Moon, Mars, EPB or also study their analogs)

- Chemistry & mineralogy of vaporization residue as a function of the fraction vaporized
- 2. Partial pressures of SiO, O₂, Na, K, Fe, Mg as f(T)
 - a. Can compute activities of SiO₂, Na₂O, K₂O, "FeO", MgO in molten peridotite
 - b. Determine vaporization coefficients for major gases
- gives observations to test and improve MAGMA code, also guides

future astronomical observations

Effect of water vapor on vaporization chemistry

- Control H₂O partial pressures during vaporization of peridotite and constrain the importance of hydroxides/oxyhydroxides for Na, K, Fe, Mg, Si gas chemistry
- If #1 is not feasible, try olivine Fo₉₀Fa₁₀ (this is 90% Mg₂SiO₄
 Forsterite Fo 10% Fe₂SiO₄ Fayalite, Fa) and study importance of H₂O for Fe, Mg, Si gas chemistry

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